

## Characterization of Regional Earth Magnetic Activity Based On MAGDAS Data

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**Abstract**—Earth magnetic data can be measured by magnetometer, Magnetic Data Acquisition System and other devices. MAGDAS is a realtime system magnetometers that are being deployed by Kyushu Sangyo University of Fukuoka, Japan. The earth magnetic data from the MAGDAS can be used as part of earth magnetic data monitoring, mainly for earth events such as earthquake and space events such as geomagnetic storm. For this project, data from stations at the north, equatorial and south region are considered as monitored data to be analyzed. The analyses are based on the different biases of MAGDAS data at different region. From the analysis, the earth magnetic polar at different region can be mapped and defined. The results are useful to give the regional of magnetic format data.

**Keywords**— Magnetic Data Acquisition System (MAGDAS), Global Positioning System (GPS), Planetary index ( $Kp$  index)

### 1.0 INTRODUCTION

Magnetic Data Acquisition System (MAGDAS) is a system of 50 realtime magnetometers that are being deployed by Kyushu Sangyo University of Fukuoka, Japan. On the April 2007 the deployment was concentrated along the 210 magnetic meridian. However, during the current stage of expansion, units are also being deployed along the geomagnetic equator, in places such as Malaysia, Ethiopia, Nigeria, Ivory Coast, and also Brazil. The overview of study is this project will focus on evaluating MAGDAS data at the north, centre and south region of the world. We chose to study this because we want to identify study and evaluate the different earth magnetic field located at the north, centre and south stations of the world. The problem statement for this project is there is no clear characterization of the earth magnetic activity based on MAGDAS data. Therefore, this project is focus on the characterization of earth magnetic field. Data are taken from the north, centre and south of the world to identify the characterization of the earth magnetic field.

### 2.0 MAGDAS DATA AND EARTH MAGNETIC FIELD

Magnetic field digital data such as  $H+\delta H$ ,  $D+\delta D$ ,  $Z+\delta Z$ ,  $F+\delta F$  are obtained at the sampling rate of 1/16 seconds, and then the averaged data are transferred from overseas

stations to SERC. The ambient magnetic field which is expressed by horizontal (H), declination (D), and vertical (Z) components, are digitized by using the field-canceling coils for the dynamic range of  $\pm 64,000\text{nT}/16\text{bit}$ . Next, the total field ( $F+\delta F$ ) is estimated from the  $H+\delta H$ ,  $D+\delta D$ , and  $Z+\delta Z$  components. For the resolution, resolution of MAGDAS data are 0.061 nT/LSB and 0.031nT/LSB for  $\pm 2,000$  nT and  $\pm 1,000$  nT range. For further information, the estimated noise level of the MAGDAS magnetometers is 0.02nTp-p, the long-term inclinations denoted as I of the sensor axes are measured by two tiltmeters with 0.2 arc-sec resolution. The temperature, denoted as T inside the sensor head is also measured. The GPS signals are received to adjust the standard time inside the data logger/transfer unit. These data are logged into the Compact Flash Memory Card and the total weight of the compact MAGDAS magnetometer system is less than 15 kg. Below is the calculation on the earth magnetic field. The x, y and z are the coordinate.

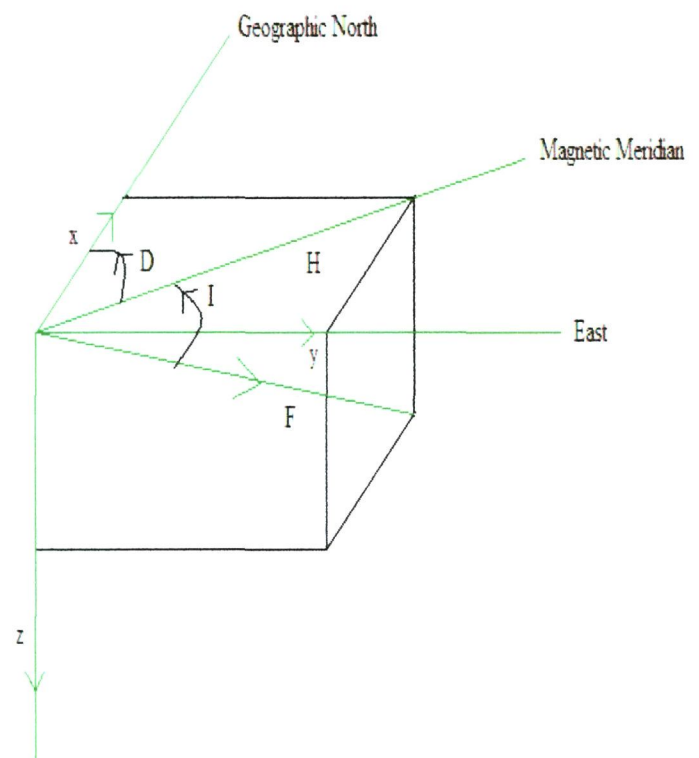


Fig. 1. Earth Magnetic Field

To get F,

$$F = \sqrt{x^2 + y^2 + z^2} \quad (1)$$

To get H,

$$H = \sqrt{x^2 + y^2} \quad (2)$$

To get angle I,

$$I = \tan^{-1} z/H \quad (3)$$

To get angle D,

$$D = \cos^{-1} x/H \quad (4)$$

The definition of the MAGDAS parameter is explained here. The declination (D) indicates the difference, in degrees, between the headings of true north and magnetic north. For inclination, inclination (I) is the angle, in degrees, of the magnetic field above or below horizontal. Next, horizontal Intensity (H) defines the horizontal component of the total field intensity and vertical Intensity (Z) defines the vertical component of the total field intensity. Last is the total Intensity (F) is the strength of the magnetic field.

### 3.0 METHODOLOGY

In specific methodology, the first process is to select the station to be analyzed. The station to be analyzed can be at Langkawi, Malaysia, Onagawa, Japan and Cooktown, Australia. After selecting the station, MAGDAS data is extract from MGD file to text document and also save in folder work. After selecting the station, insert the file name for example ONW\_MIN\_200609090000.mgd into the read\_1M.m. Next, if the data is not available in the command window, the MAGDAS data is extract again from MGD file to text document and read by using read\_1M.m again. If the data is available, the data is shown in the command window. Afterward, choose which bias to be plotted and don't forget to insert the time in the workspace. The time can be in 24 hour depending on the location. Next, click New variable in the Workspace, rename it with a suitable name and plot the bias with time. Type, for example plot (time,Bias\_H) on the command window. The graph is then plotted and analyze.

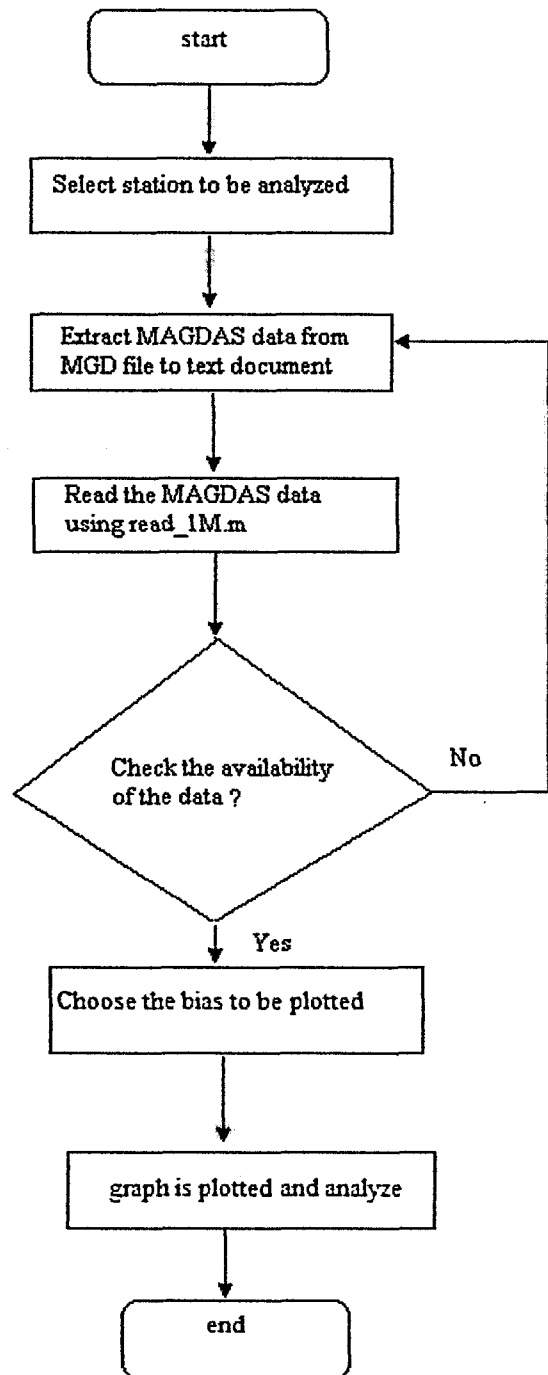


Fig. 2. Methodology

**TABLE 1**  
**THE CHARACTERIZATION OF EARTH**  
**MAGNETIC ACTIVITY BASED ON DATE AND**  
**LOCATION**

| Region     | Station             | Date   |
|------------|---------------------|--|
| Northern   | Onagawa, Japan      | 9 <sup>th</sup> September 2006 and 10 <sup>th</sup> September 2006 |
| Equatorial | Langkawi, Malaysia  | 9 <sup>th</sup> September 2006 and 10 <sup>th</sup> September 2006 |
| Southern   | Cooktown, Australia | 9 <sup>th</sup> September 2006 and 10 <sup>th</sup> September 2006 |

**4.0 RESULTS**

The results show bias\_F at Onagawa, Japan, Langkawi, Malaysia and Cooktown, Australia. Bias\_F is use to show the strength of the magnetic field.

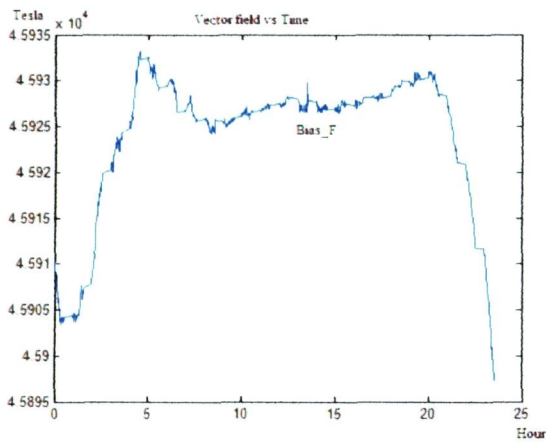


Fig 3. Bias\_F for Onagawa, Japan on 9<sup>th</sup> September 2006

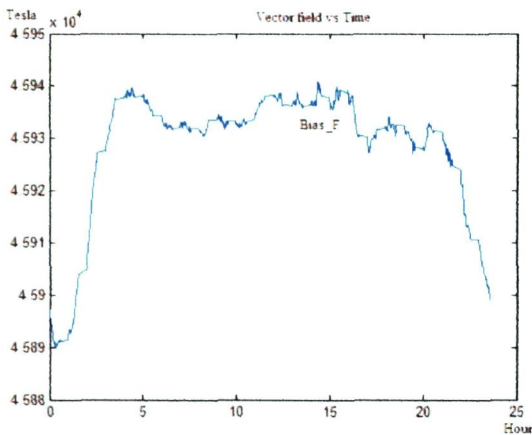


Fig 4. Bias\_F for Onagawa, Japan on 10<sup>th</sup> September 2006

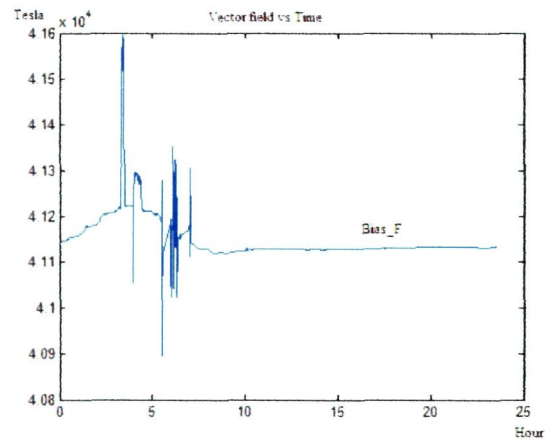


Fig 5. Bias\_F for Langkawi, Malaysia on 9<sup>th</sup> September 2006

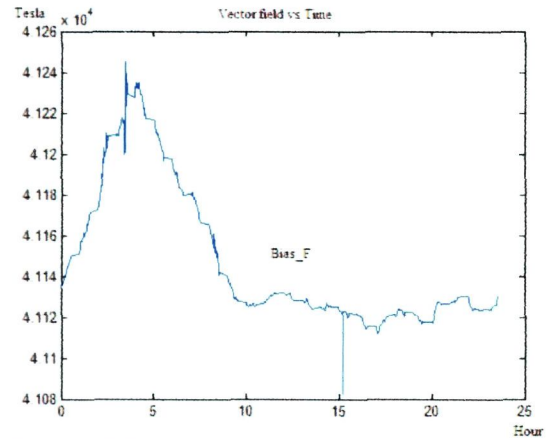


Fig 6. Bias\_F for Langkawi, Malaysia on the 10<sup>th</sup> September 2006

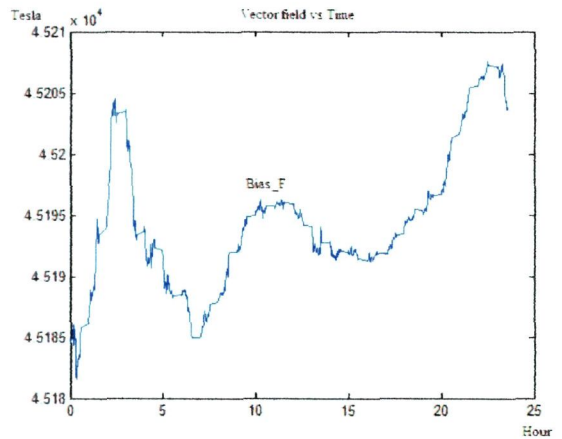


Fig 7. Bias\_F for Cooktown, Australia on the 9<sup>th</sup> September 2006

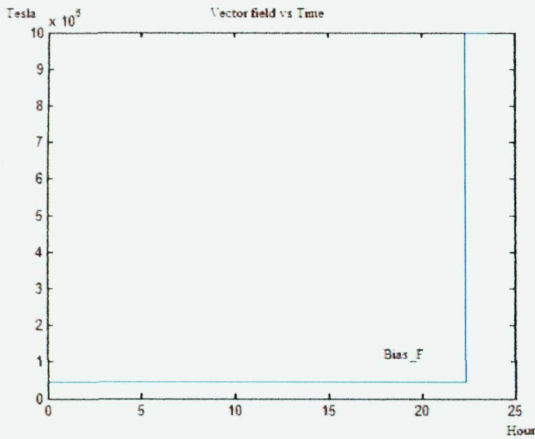


Fig 8. Bias\_F for Cooktown, Australia on the 10 September 2006

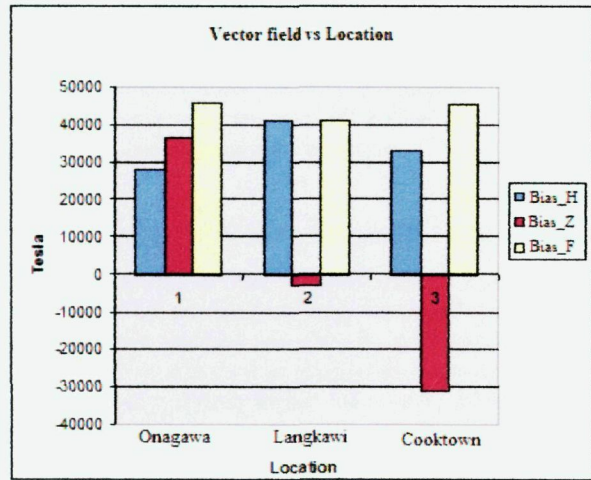


Fig 9. The average data for bias\_H, bias\_Z and bias\_F on 9th September 2006

TABLE 2  
THE AVERAGE DATA FOR EARTH MAGNETIC FIELD ON 9<sup>th</sup> SEPTEMBER 2006

| Station             | The average data for bias H | The average data for bias D | The average data for bias Z | The average data for bias F |
|---------------------|-----------------------------|-----------------------------|-----------------------------|-----------------------------|
| Onagawa, Japan      | 28005 T                     | -7.81°                      | 36405 T                     | 45916.60 T                  |
| Langkawi, Malaysia  | 41100 T                     | -145°                       | -2802.77 T                  | 41150 T                     |
| Cooktown, Australia | 32985 T                     | -7.81°                      | -30905 T                    | 45201.75 T                  |

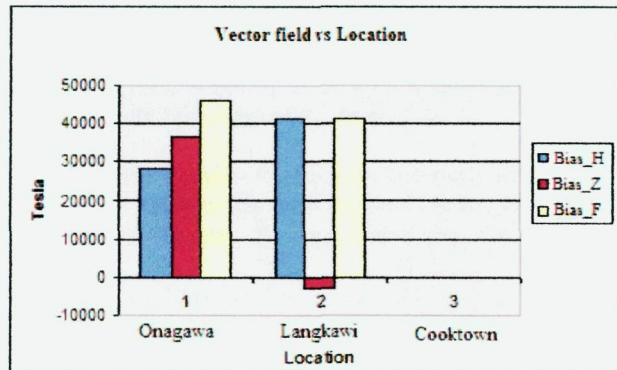


Fig 10. The average data for bias\_H, bias\_Z and bias\_F on 10th September 2006

TABLE 3  
THE AVERAGE DATA FOR EARTH MAGNETIC FIELD ON 10 SEPTEMBER 2006

| Station             | The average data for bias H | The average data for bias D | The average data for bias Z | The average data for bias F |
|---------------------|-----------------------------|-----------------------------|-----------------------------|-----------------------------|
| Onagawa, Japan      | 28005 T                     | -7.81°                      | 36410 T                     | 45916.60 T                  |
| Langkawi, Malaysia  | 41050 T                     | 3.0°                        | -2802.77 T                  | 41130 T                     |
| Cooktown, Australia | No Measurement              | No Measurement              | No Measurement              | No Measurement              |

### 5.0 DISCUSSION

The Earth's magnetic field is generated within its molten iron core through a combination of thermal movement, the Earth's daily rotation, and electrical forces within the core. These elements form a dynamo that sustains a magnetic field that is similar to that of a bar magnet slightly inclined to a line that joins the North and South Geographic Poles. Based on the result, we can see the characterization of the earth magnetic field. The total intensity (F) which shows the strength of magnetic field is larger at the north and south of the world compare to the center of the world since most of the earth magnetic activity concentrate at the North and South Geographic Poles. Table 1 and table 2 shows the average data for the earth magnetic field on 9<sup>th</sup> and 10<sup>th</sup> September 2006. It shows the average data for bias\_F is larger at the north and south compare to the center of the world.

On the 9<sup>th</sup> to 10<sup>th</sup> September 2006, the k-Index shows there is no geomagnetic storm occurs since the value of the k-index is not greater than 5.

After researching, the most important bias is bias<sub>H</sub> since bias<sub>H</sub> can be used to detect magnetic storm. In the website <http://swdcwww.kugi.kyoto-u.ac.jp/dst/dir/dst2/onDstindex.html>, it states 'the onset of a magnetic storm is often characterized by a global sudden increase in H, which is referred to as the storm sudden commencement and denoted by ssc. Following the ssc, the H component typically remains above its average level for a few hours; this phase is called the initial phase of the storm. Then a large global decrease in H begins, indicating the development of the main phase of the storm. The magnitude of the decrease in H represents the severity of disturbance'.

The parameter of the earth magnetic field for example horizontal intensity (H) is most directly influenced by the external field contributions. Variations in vertical intensity (Z) are strongly influenced by the electrical conductivity of the subsurface and contain information on the field contributions induced in the Earth by the external variations.

The graph for Bias<sub>H</sub>, figure 14 for Cooktown, Australia on the 10<sup>th</sup> September 2006 shows a horizontal line and a vertical line. On that day, there was no measurement being made which explains why the graph was a horizontal line and vertical line.

The geomagnetic variations are usually caused by 3 factors which are core field variations, solar variations and solar wind-magnetosphere interaction. For Bias<sub>F</sub> at Langkawi, Malaysia, Onagawa, Japan and Cooktown, Australia from figure 3 to figure 8 are affected by solar radiation and core field variation. Solar radiations are changes in the amount of solar radiation emitted by the sun. Interaction of solar particles, the solar magnetic field, and the earth's magnetic field, cause variations in the particle and electromagnetic fields at the surface of the planet. The peak in the geomagnetic variation for figure 5 is caused by core field variation which consists of reverse, westward drift and secular variation.

## 6.0 CONCLUSION

For the conclusion, the characterization of the earth magnetic field has been seen. The earth magnetic field is greater at the north and south compared to the center of the world. The study in the characterization of the earth magnetic activity using MAGDAS data has made us realize the character of earth magnetic field at the north, center and south of the world. In the end, this project is important since the study in the characterization of the earth

magnetic field can provide information on the earth magnetic field located at the south, north and center region of the world. Besides, the variation of earth magnetic field could detect major events such as geomagnetic storm. The information gain can be used to minimize the cost of destruction caused by geomagnetic storm.

## 7.0 FUTURE RECOMMENDATION

Future research should be conducted to cover a more detail study about this project. They are as follows, take more data at different stations and analyze it, research based on the other earth events such as earthquake, volcano, flood and seasonal events and study higher value of KP-index to detect geomagnetic storm.

## REFERENCES

- [1] Faizatul Noor binti Abu Bakar, "Determination of traveling ionospheric disturbances (TID) of geomagnetic storm by using dual frequency GPS data", Universiti Teknologi Mara, May 2008.
- [2] Kiyohumi Yumoto, "Studies on Geomagnetic Field and the Relationship with the Sun", Space Environment Research Center, Kyushu University, 26 December 2006.
- [3] Aoi Nakamizo, "Manual for Space Weather Nowcasting", Space Environment Research Center, Kyushu University, 15 November 2006.
- [4] Micheal Browne, "Physics for engineering and science", Mc Graw Hill, 1999
- [5] Frederick J. Bueche, "College Physics", Mc Graw Hill, 1997
- [6] Bias<sub>H</sub>, available (March 2009): <http://swdcwww.kugi.kyoto-u.ac.jp/dst/dir/dst2/onDstindex.html>